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Vegetation dynamics and human activity in the Western Pyrenean Region during the Holocene

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ABSTRACT

The present paper summarises the evolution of the vegetation and the evidences of anthropisation in the Western Pyrenees (Northern Iberian Peninsula) during the Holocene. All the palynological studies published about this region, about both natural and archaeological deposits, have been compiled, and special attention is paid on the available radiocarbon dates. We also present new results of the palynological study of the peat bog of Atxuri (Navarre). The main results document the arboreal colonisation in the early Holocene; the first evidences of anthropisation associated with the Early Neolithic (ca. 5500–4500 cal BC); the consolidation of a productive economy (agriculture and pastoralism) in the Middle/Late Neolithic (ca. 4500–3200 cal BC) and above all in the Chalcolithic (ca. 3200–2200 cal BC); and the rise of complex urban societies in the Bronze Age (ca. 2200–900 cal BC) and Iron Age (ca. 900–200 cal BC) and their impact on the vegetation.

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1. Introduction

The Western Pyrenees is a region whose climate, geology, biogeography and history have resulted in a great environmental heterogeneity and biological diversity, with a noteworthy wealth of species, biological communities and ecosystems. In addition, from the viewpoint of its geographical situation, it has been on the traditional route from central and southern Europe to the inner Iberian Peninsula, as shown by the large number of archaeological remains of different periods that have been found (Barandiarán-Maestu et al., 1998; Alday et al., 2006; Peñalver, 2008; Fernández-Eraso et al., 2009).

The current vegetation of this region has originated in different processes that have taken place over time and which have modified the structure of the forests significantly (Costa-Tenorio et al., 2005). These phenomena correspond, on one hand, to continuous environmental changes and, on the other, to the weight of a long heterogeneous human use of the territory. The latter factor, the impact of human communities on the environment, becomes particularly noticeable after ca. 5500 cal BC, mainly owing to the change in the

ways of life of the prehistoric communities. After this time, farming practices spread across the Western Pyrenees, significantly modifying local ecosystems in both mountain and valley areas, within a process related to the Neolithisation dynamic in the Mediterranean basin (Galop et al., 2013).

To obtain an accurate picture of the phenomena involved in shaping the modern vegetation communities it is necessary to take into account all the available palaeobotanical records, especially palynological studies both from archaeological sediments and from natural deposits like peat bogs, lakes, estuaries, and so on. All these data together provide the essential information to address some issues of great interest and hot debated, such as the evolution of the main plant communities over time, the appearance/expansion/regression of some taxa of particular interest because of their rareness, climate change and its connection with the development of the vegetation, relationships between human communities and the environment, and aspects such as the anthropisation processes associated with the Neolithic (Dupré, 1988; Birks et al., 1988; Berglund, 1991; Galop, 1998; López-Sáez et al., 2000, 2003; Barbier et al., 2001; Davis et al., 2003; Carrión et al., 2010; Vannière et al., 2011; Nieto-Moreno et al., 2011; Galop et al., 2013; Sadori et al., 2013; among others).

The first palynological research in the region took place in the 1980s when this type of study began to be widely applied both to natural deposits (Peñalba, 1989) and to archaeological ones (López-

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García, 1982; Dupré, 1984; Boyer-Klein, 1985; García-Antón et al., 1987; Peñalba, 1987; Sánchez-Goni, 1987). Since then, more than 40 new palynological sequences have been published both in Spanish journals (Iriarte, 2006, 2009b; Pérez-Díaz and López-Sáez, 2013) and International journals (Sanchez-Goni, 1996; Iriarte, 2009a; Corella et al., 2013).

This paper focuses on the evolution of the vegetation over the last 9500 years, in order to evaluate the anthropic impact on the landscape by prehistoric groups that inhabited the Western Pyrenees. With this objective, all the available palynological studies from the Epipaleolithic-Mesolithic to the Iron Age, and a full series of 219 radiocarbon dates, have been taken into account.

2. Regional setting

The Western Pyrenees is quite rugged, with great climatic and geomorphologic variety. The relief is generally aligned east-west, and therefore the area consists of a series of mountain barriers, parallel to the coastline, which limits the spread of Atlantic flora southwards. From the biogeographic point of view, it can be divided into three areas: the Atlantic zone (coastal and inner valleys of the north of the Basque Country), the Pyrenees (north of Navarre), both belonging to the Eurosiberian region, and the Transition area (Sub-Atlantic and Sub-Mediterranean Valleys of Álava), belonging to the Mediterranean region (Fig. 1).

The Pyrenees is a large mountain range over 400 km long, separating the Aquitaine basin from the Ebro valley. Its main valleys in the north of Navarre run perpendicular to the range. The climate is generally cold and wet in winter, when snow and mist are common. In contrast, summer temperatures can be relatively high. It has large well-conserved forest masses with *Pinus sylvestris* L. (Scots pine), *Quercus robur* L. (Pendunculate oak), *Quercus faginea*

Lam. (Portuguese oak), *Fagus sylvatica* L. (beech) and *Abies alba* Mill. (fir).

In northern Iberia, the Atlantic zone is a strip parallel to the coastline, directly influenced by the sea. It is characterised by undulating relief, with hills and wide valleys near the coast and more abrupt relief inland. The climate is very humid and temperate, with hardly any summer drought (Aseginolaza et al., 1996). The main elements in the vegetation today are anthropic, with farmlands, above all meadows, together with plantations of conifers. Stands of *Quercus ilex* L. (holm oak) are found in some limestone areas, and small dispersed woods with pendunculate oak and other deciduous species (Aseginolaza et al., 1996).

Finally, the climate of the Transition area displays characteristics of both Atlantic and Mediterranean environments. In high mountain areas the climate is characterised by low temperatures and heavy precipitation, with frequent mist, and frost and snow except in summer. The climatic and geomorphologic conditions result in vegetation communities consisting of calcicolous beech woods on limestone hills, together with anthropic pastures used by livestock. In the valleys, the most important vegetation types are oak woods with pendunculate oak, *Quercus pyrenaica* Willd. (Pyrenean oak) and small stands of Portuguese oak, together with areas of holm oak and Scots pine (Aseginolaza et al., 1996).

3. Materials and methods

All palynological studies published for the Holocene on the southern side of the Western Pyrenees have been examined. In total, these refer to 55 deposits, of which 7 are natural deposits and 48 are archaeological sites. They are all situated chronologically with a total of 219 absolute dates (Table 1).

Table 1

Holocene pollen sequences in the study area. Nature: AS (Archaeological Site) – Mg (Megalithic Monument), C (Cave), OA (Open Air Site), RS (Rock Shelter). ND (Natural Deposit) – M (Marsh), P (Peatbog), L (Lake). Radiocarbon dates have been calibrated with the Calib 7.0 programme at 2σ (95.4% probability) and are given in dates BP (in brackets cal BC or cal AD).

Number	Site	Location	Nature	Altitude	Age BP (Age cal BC/AD)	References
Atlantic area						
1	Aitxu	Ataun-Idiazabal	AS (Mg)	930	3530 ± 110 (2196–1566 cal BC)	Iriarte, 1997c
2	Amalda	Cestona	AS (C)	205	1740 ± 200 (185 cal BC–665 cal AD), 1460 ± 80 (415–762 cal AD)	Dupré, 1990
3	Arenaza I	Galdames	AS (C)	185	10,300 ± 180 (10,653–9419 cal BC), 9600 ± 180 (9446–8354 cal BC), 6040 ± 75 (5207–4779 cal BC), 5755 ± 65 (4767–4457 cal BC), 4965 ± 195 (4259–3346 cal BC), 4730 ± 110 (3760–3108 cal BC), 3835 ± 55 (2466–2141 cal BC), 3805 ± 70 (2465–2039 cal BC), 3580 ± 70 (2135–1746 cal BC)	Isturiz and Sanchez-Goni, 1990
4	Berreaga	Mungia	AS (OA)	360	–	Iriarte, 1994a
5	Buruntza	Andoain	AS (OA)	439	3000 ± 60 (1401–1055 cal BC), 2810 ± 90 (1253–806 cal BC), 2475 ± 75 (779–408 cal BC), 2270 ± 80 (706–95 cal BC), 2180 ± 80 (394–46 cal BC)	Iriarte, 1997b
6	Ekain	Deba	AS (C)	90	9460 ± 185 (9249–8323 cal BC), 9540 ± 210 (9433–8295 cal BC), 9610 ± 85 (9247–8766 cal BC), 4960 ± 60 (3942–3641 cal BC), 5179 ± 170 (4338–3653 cal BC), 4680 ± 60 (3633–3358 cal BC), 4120 ± 50 (2876–2505 cal BC), 3820 ± 240 (2896–1658 cal BC), 3700 ± 40 (2202–1973 cal BC)	Dupré, 1984
7	Gastiburu	Arratzu	AS (OA)	340	3260 ± 60 (1681–1427 cal BC), 2455 ± 26 (753–412 cal BC), 2396 ± 27 (723–398 cal BC), 2270 ± 26 (398–211 cal BC), 2257 ± 29 (394–209 cal BC), 2233 ± 26 (386–205 cal BC), 2190 ± 26 (362–180 cal BC), 2157 ± 27 (357–106 cal BC), 2140 ± 60 (370–41 cal BC), 2076 ± 44 (201 cal BC–21 cal AD), 2060 ± 60 (345 cal BC–69 cal AD), 1759 ± 28 (176–383 cal AD)	Valdés, 2009
8	Herriko Barra	Zarauz	AS (OA)	5	–	Iriarte et al., 2004

Table 1 (continued)

Number	Site	Location	Nature	Altitude	Age BP (Age cal BC/AD)	References
					5810 ± 170 (5203–4332 cal BC), 6010 ± 90 (5207–4713 cal BC), 5960 ± 95 (5204–4590 cal BC)	
					5730 ± 110 (4824–4352 cal BC), 5800 ± 110 (4933–4374 cal BC), 4920 ± 100 (3959–3390 cal BC)	
9	Ilso Betaio	Arcentales-Sopuerta	AS (OA)	712	930 ± 80 BP (981–1261 cal AD)	Iriarte, 1999a
10	Intxur	Albiztur-Tolosa	AS (OA)	742	2720 ± 80 (1112–769 cal BC), 2400 ± 80 (782–265 cal BC), 2260 ± 80 (519–60 cal BC)	Iriarte, 1994a
					2180 ± 80 (394–46 cal BC), 2070 ± 80 (358 cal BC–80 cal AD), 2030 ± 80 (352 cal BC–132 cal AD)	
11	Katillotxu	Mundaka	AS (Mg)	336	5200 ± 40 (4225–3948 cal BC), 4290 ± 40 (3022–2780 cal BC)	Iriarte, 2007/2009
12	Kobaederra	Kortezubi	AS (C)	260	5630 ± 100 (4709–4270 cal BC), 5375 ± 90 (4359–3989 cal BC), 5820 ± 240 (5306–4175 cal BC)	Iriarte, 2009a
					5460 ± 90 (4460–4050 cal BC), 5200 ± 110 (4323–3769 cal BC), 4405 ± 55 (3331–2905 cal BC)	
13	Kobeaga II	Ispster	AS (C)	205	7690 ± 270 (7295–6018 cal BC), 6945 ± 65 (5983–5718 cal BC)	Iriarte, 2000
14	Kosnoaga	Gernika	AS (OA)	274	–	Iriarte, 1994a
15	La Boheriza 2	Carranza	AS (Mg)	732	5500 ± 100 (4542–4054 cal BC), 5200 ± 75 (4235–3802 cal BC)	Iriarte, 1995a
16	La Cabaña 4	Carranza	AS (Mg)	713	–	Iriarte, 1999c
17	Laminak II	Berriatua	AS (C)	40	11,700 ± 140 (11,867–11,328 cal BC), 10,380 ± 140 (10,654–9769 cal BC)	Isturiz, 1994
18	Lezetxiki	Mondragón	AS (C)	380	–	Sanchez-Goni, 1993
19	Mulixko Gaina	Urnieta-Hernani	AS (Mg)	415	2630 ± 90 (1002–420 cal BC)	Peñalba, 1987
20	Napalatz	Idiazabal	AS (Mg)	740	1830 ± 90 (1–470 cal AD), 1670 ± 140 (77–642 cal AD)	Iriarte, 1999b
21	Pareko Landa	Bermeo-Busturia	AS (OA)	522	7510 ± 100 (6590–6107 cal BC), 6650 ± 130 (5793–5330 cal BC)	Iriarte et al., 2007/2008
22	Pico Ramos	Muskiz	AS (C)	215	6840 ± 75 (5890–5622 cal BC), 6040 ± 90 (5210–4729 cal BC), 5860 ± 65 (4895–4547 cal BC), 5370 ± 40 (4331–4055 cal BC)	Iriarte, 1994b
					4790 ± 110 (3910–3199 cal BC), 4210 ± 110 (3090–2480 cal BC), 4100 ± 110 (2910–2347 cal BC)	
23	Playaundi	Irún	ND (M)	2	7810 ± 130 (7044–6451 cal BC), 6630 ± 120 (5741–5331 cal BC), 6590 ± 120 (5716–5324 cal BC), 2740 ± 90 (1192–676 cal BC)	Sánchez Goni, 1996
24	Praalata	Ataun-Idiazabal	AS (Mg)	931	4470 ± 70 (3357–2929 cal BC), 4310 ± 110 (3338–2624 cal BC)	Iriarte, 1997c
25	Saldropo	Zeanuri	ND (P)	625	5630 ± 70 (4653–4340 cal BC), 3590 ± 90 (2199–1694 cal BC), 840 ± 50 (1045–1273 cal AD)	García-Antón et al., 1987; Peñalba, 1989
					4510 ± 140 (3629–2891 cal BC), 2460 ± 100 (803–387 cal BC), 920 ± 100 (901–1277 cal AD)	
26	Sta. María la Real	Zarauz	AS (OA)	7	2370 ± 40 (733–382 cal BC)	Iriarte, 2009b
27	Santimamiñe	Kortezubi	AS (C)	137	10,100 ± 60 (10,080–9440 cal BC), 10,060 ± 60 (10,030–9360 cal BC), 7580 ± 50 (6490–6380 cal BC)	Iriarte, 2011
					5450 ± 50 (4360–4230 cal BC), 5010 ± 40 (3940–3700 cal BC), 3710 ± 40 (2269–1977 cal BC)	
28	Sustrigi	Areatza-Bilaro	AS (OA)	734	–	Iriarte et al., 2007/2008; Iriarte, 2009a
29	Unanibi	Idiazabal	AS (Mg)	840	3350 ± 100 (1886–1433 cal BC), 3085 ± 75 (1500–1128 cal BC), 3040 ± 230 (1878–789 cal BC)	Iriarte, 1999b
30	Urdaibai	Gernika	ND (M)	12	9050 ± 40 BP (8303–8228 cal BC), 2580 ± 40 (820–549 cal BC)	Iriarte et al., 2006
31	Urtiaga	Deba	AS (C)	160	8700 ± 170 (8259–7484 cal BC)	Sanchez-Goni, 1993
32	Zorroztarri	Idiazabal-Segura	AS (Mg)	678	3280 ± 90 (1865–1326 cal BC), 3040 ± 90 (1493–1024 cal BC)	Iriarte, 1994a
Pyrenean area						
33	Abauntz	Arraiz	AS (C)	709	9530 ± 300 (9976–7992 cal BC), 6910 ± 450 (6247–4796 cal BC), 5820 ± 40 (4781–4556 cal BC)	López-García, 1982
					5390 ± 120 (4452–3973 cal BC), 4240 ± 140 (3332–2474 cal BC), 4370 ± 70 (3332–2883 cal BC)	
					4025 ± 35 (2829–2469 cal BC), 3975 ± 35 (2577–2349 cal BC), 3900 ± 35 (2475–2244 cal BC)	
34	Aizpea	Aribe	AS (RS)	720	7790 ± 90 (7024–6454 cal BC), 7160 ± 70 (6211–5902 cal BC), 6830 ± 70 (5876–5622 cal BC)	Iriarte, 2001
					6600 ± 50 (5620–5482 cal BC), 6370 ± 70 (5475–5222 cal BC)	
35	Atxuri	Maya de Baztán	ND (P)	500	2710 ± 60 (995–796 cal BC), 1920 ± 80 (109 cal BC–322 cal AD)	Peñalba, 1989; Galop et al., 2004
					6370 ± 80 (5491–5081 cal BC), 3880 ± 40 (2470–2210 cal BC), 3300 ± 40 (1683–1499 cal BC), 2180 ± 40 (371–113 cal BC)	

(continued on next page)

Table 1 (continued)

Number	Site	Location	Nature	Altitude	Age BP (Age cal BC/AD)	References
36	Belate	Narvarte	ND (P)	847	6600 ± 80 (5666–5380 cal BC), 5900 ± 80 (4962–4553 cal BC), 2960 ± 60 (1386–1004 cal BC)	Peñalba, 1989
37	Berroberría	Urdax	AS (C)	110	9740 ± 140 (9660–8735 cal BC), 8860 ± 100 (8263–7676 cal BC), 8800 ± 80 (8212–7612 cal BC), 8630 ± 70 (7935–7536 cal BC), 8580 ± 80 (7816–7484 cal BC), 8510 ± 90 (7731–7352 cal BC), 8470 ± 80 (7610–7329 cal BC), 8130 ± 90 (7451–6780 cal BC), 7640 ± 90 (6652–6265 cal BC), 2655 ± 35 (985–793 cal BC)	Boyer-Klein, 1984
38	Padre Areso	Biguezabal	AS (OA)	900	5400 ± 100 (4447–3993 cal BC), 3020 ± 35 (1394–1129 cal BC), 1130 ± 60 (731–1019 cal AD)	Iriarte et al., 2007/2008
39	Quinto Real	Valle de Baigorri	ND (P)	910	4120 ± 40 (2871–2577 cal BC), 3045 ± 70 (1449–1058 cal BC), 2645 ± 45 (901–775 cal BC), 1895 ± 50 (5–237 cal AD), 1155 ± 35 (775–969 cal AD), 290 ± 40 (1784–1795 cal AD)	Galop et al., 2004
40	Zatoya	Abaurrea Alta	AS (C)	900	8260 ± 550 BP (8618–6033 cal BC), 8150 ± 220 (7584–6595 cal BC), 6320 ± 280 (5743–4590 cal BC)	Boyer-Klein, 1989
Transition area						
41	Amezutxo	Ilarduia	AS (OA)	594	–	Pérez-Díaz 2012
42	Arreo	Lantarón	ND (L)	655	2480 ± 70 (782–414 cal BC), 1860 ± 40 (65–243 cal AD), 1550 ± 35 (420–584 cal AD), 1400 ± 60 (541–767 cal AD), 1170 ± 70 (688–992 cal AD)	Corella et al., 2013
43	Atxa	Vitoria-Gasteiz	AS (OA)	525	2410 ± 90 (792–261 cal BC), 2320 ± 90 (755–177 cal BC), 2280 ± 80 (730–111 cal BC), 2200 ± 80 (397–54 cal BC), 2180 ± 80 (394–46 cal BC)	Iriarte, 1994a, 1995b
44	Berniollo	Subijana-Morillas	AS (OA)	520	9940 ± 490 (10,805–8229 cal BC), 4160 ± 190 (3337–2206 cal BC), 3910 ± 100 (2836–2044 cal BC)	Isturiz and Sanchez-Goni, 1990
45	Kanpanoste	Vírgala	AS (RS)	700	8200 ± 70 (7451–7059 cal BC), 7920 ± 100 (7072–6534 cal BC), 7620 ± 70 (6634–6371 cal BC)	Sánchez-Goni, 2004
46	Kanpanoste Goikoa	Vírgala	AS (RS)	720	7860 ± 330 (7528–6095 cal BC), 7620 ± 80 (6641–6266 cal BC), 6550 ± 260 (5986–4913 cal BC), 6360 ± 70 BP (5475–5217 cal BC), 4550 ± 40 (3485–3100 cal BC), 4350 ± 60 (3321–2878 cal BC), 4190 ± 100 (3013–2490 cal BC)	Iriarte, 1998
47	Kukuma	Araia	AS (C)	715	11,550 ± 130 (11,762–11,215 cal BC)	Isturiz, 1997
48	La Peña	Marañón	AS (RS)	640	7890 ± 120 (7065–6491 cal BC), 4350 ± 80 (3338–2765 cal BC), 3710 ± 60 (2286–1943 cal BC), 3610 ± 60 (2140–1775 cal BC), 2840 ± 70 (1207–840 cal BC)	López-García, 1991/1992
49	Los Castros de Lastra	Caranca	AS (OA)	830	2990 ± 80 (1419–1005 cal BC), 2800 ± 90 (1249–802 cal BC), 2760 ± 90 (1193–787 cal BC), 2750 ± 90 (1192–774 cal BC), 2730 ± 170 (1313–414 cal BC), 2720 ± 80 (1112–769 cal BC), 2680 ± 100 (1116–538 cal BC), 2630 ± 190 (1300–259 cal BC), 2580 ± 80 (897–417 cal BC), 2530 ± 80 (806–413 cal BC), 2520 ± 80 (808–407 cal BC), 2300 ± 80 (749–577 cal BC), 2140 ± 80 (383–0)	Pérez-Díaz, 2012
50	Masparra	Margarita	AS (OA)	504	–	Pérez-Díaz and López-Sáez, 2013
51	Mendandia	Sáseta	AS (RS)	720	8500 ± 60 (7605–7460 cal BC), 7810 ± 50 (6803–6500 cal BC), 7780 ± 60 (6754–6468 cal BC), 7620 ± 50 (6590–6410 cal BC), 7210 ± 45 (6211–6004 cal BC), 7180 ± 45 (6206–5933 cal BC), 6540 ± 70 (5620–5371 cal BC), 6440 ± 40 (5478–5330 cal BC), 2860 ± 35 (1187–918 cal BC)	Iriarte, 2006
52	Mendizabala	Vitoria-Gasteiz	AS (OA)	534	6720 ± 40 (5715–5560 cal BC), 6150 ± 230 (5524–4544 cal BC), 5830 ± 110 (4977–4451 cal BC)	Pérez-Díaz 2012
53	Peña Larga	Cripán	AS (RS)	900	5720 ± 40 (4684–4463 cal BC), 5010 ± 40 (3944–3704 cal BC), 4890 ± 50 (3787–3536 cal BC), 4550 ± 40 (3485–3100 cal BC), 4470 ± 160 (3633–2702 cal BC), 4040 ± 40 (2838–2469 cal BC)	Iriarte, 1997a, Pérez-Díaz 2012
54	Peña Parda	Laguardia	AS (RS)	975	–	Pérez-Díaz et al., 2007
55	San Cristóbal	Laguardia	AS (RS)	1037	5100 ± 30 (3968–3800 cal BC), 4980 ± 30 (3911–3663 cal BC), 4930 ± 30 (3771–3651 cal BC), 4470 ± 40 (3345–3022 cal BC), 4290 ± 40 (3022–2780 cal BC), 4170 ± 40 (2886–2628 cal BC), 4100 ± 40 (2870–2498 cal BC), 4080 ± 40 (2862–2489 cal BC), 4030 ± 40 (2835–2468 cal BC), 3990 ± 40 (2620–2350 cal BC)	Pérez-Díaz, 2012

Iversen (1989), although some identification required the use of a pollen atlas (Reille, 1992; Beug, 2004). The palynological diagram (Fig. 2) was produced using the Tilia 2.0 and TGView programs (Grimm, 1992, 2004).

Table 2
AMS radiocarbon dates from the peat bog of Atxuri (Navarre). Radiocarbon dates have been calibrated with the Calib 7.0 programme at 2σ (95.4% probability) and are given in cal BC.

Sample	Material	Depth interval	Lab code	Age uncalibrated BP	Age cal BC (2σ range probability area)	Medium age cal BC
Atxuri-55	Peat	54–55	BETA-156993	2180 ± 40	371–113	265
Atxuri-130	Peat	129–130	BETA-156995	3300 ± 49	1683–1499	1577
Atxuri-160	Peat	159–160	BETA-156994	3880 ± 80	2470–2210	2368
Atxuri-230	Peat	229–230	GIFA-99149	6370 ± 80	5491–5081	5355

4. Results and discussion

4.1. Pollen zones of the Atxuri pollen diagram

The new pollen percentage diagram of the Atxuri peat bog was divided into five main biostratigraphic zones (Fig. 2) using CONISS, and provides insight into the vegetation history and the human pressure of the Pyrenean region from ca. 5500 cal BC to the present day. The first period (PAZ ATX-1, 235–185 cm depth, ca. 5568–3434 cal BC) is dominated by mesophyllous woodland, with *Corylus* and deciduous *Quercus* as main taxa, and to a lesser extent *Betula* and *Ulmus*. Among herbs, Poaceae reach highest values (20%). The PAZ ATX-2 (185–170 cm depth, ca. 3434–2795 cal BC) shows the rise of deciduous *Quercus*, *Alnus*, and *Fraxinus*, and some decline in *Corylus* and *Pinus*. Poaceae is also the main herbaceous taxa represented in the diagram. At the PAZ ATX-3 (170–135 cm depth, ca. 2795–1709 cal BC) the vegetation is dominated by mesophyllous woodland. At the beginning of this PAZ, deciduous *Quercus* continues its upward trend, reaching the maximum values of the whole sequence (39%). At this moment *Corylus* start to rise, to become again in the dominant tree taxa. The diagram shows the first appearances of *Fagus*, but in low values, the rise of *Calluna* and *Sphagnum*, and the first evidences of cultivated plants. At the PAZ ATX-4 (135–40 cm depth, ca. 1709–738 cal BC) the most important event is the spread of *Fagus*, reaching 39% and being, from this moment, the dominant woodland. Anthropogenic pollen indicators, such as *Plantago*, Urticaceae, Chenopodiaceae, and cultivated plants are more abundant at this time. Finally, the PAZ ATX-5 (40–0 cm depth, ca. 738–present day) shows a drastic drop of *Fagus*. The rest of the mesophyllous woodland also shows a downward trend. At the end of the sequence *Pinus* increases its values. Poaceae, *Calluna* and *Erica* type, together with *Pteridium*, increases significantly their presence in the vicinity of the peat bog, where anthropogenic pollen indicators and cultivated plants are also evident.

4.2. Vegetation history and human impact during the Holocene

This section is organised according to chrono-cultural periods, giving an overall view of the composition of the vegetation and the anthropic impact during the Holocene on the Spanish side of the Western Pyrenees. For this purpose we have taken into account the new data from the peat bog of Atxuri and the rest of the palynological studies published for the Pyrenees, the Atlantic area and the Transitional region, both from archaeological sites and natural ones. As we mentioned, we have considered all the studies chronologically situated between the Epipalaeolithic-Mesolithic and the Iron Age, and its radiocarbon dates. We have collected a total of 219

absolute dates. But the distribution of those radiocarbon dates is not uniform, and we have sites with more than 10 radiocarbon dates (Gastiburu, Abauntz, Berroberria, Los Castros de Lastra, Peña Larga, San Cristóbal) and others with no one (Berreaga, Kosnoaga, Sustrigi, Ameztutxo, Peña Parda). For this reason, we have been

very cautious when correlating palynological data from different sites.

4.2.1. Epipalaeolithic-Mesolithic. The environment background to the first Holocene cultures (ca. 9500–5500 cal BC)

The end of the Late Glacial and the arrival of the Holocene involved the end of Upper Palaeolithic cultures (Magdalenian), and was the starting point of new ways of life adapted to more benign climate conditions and a noticeably different flora composition. From the palaeoenvironmental viewpoint, one of the most important features of the early Holocene is a significant climatic amelioration, with a steady rise in both the temperature and precipitation (Bradley, 1999; Burroughs, 2005; Carrión et al., 2010; Muñoz-Sobrino et al., 2013).

The climate (temperature and precipitations) is, after all, one of the factors determining the vegetation communities anywhere. In the case of south-west Europe, the climate amelioration at the start of the Holocene favoured the expansion of arboreal taxa, leading to a significant reduction in the areas occupied by steppe communities, which were characteristic of the Late Glacial period (Duplessy et al., 1981; Ruddiman and McIntyre, 1981; Bard et al., 1987; Naughton et al., 2007; Fletcher et al., 2010; Jalut et al., 2010; Sadori, 2013).

Several palaeoenvironmental sequences have been documented in the Atlantic part of the region for this period. The extensive tree cover, dominated by deciduous taxa (deciduous oaks, hazel and elm) and accompanied by scots pine, is the most important trait they display. These coastal deposits also document other thermophilous elements like evergreen oak (*Q. ilex*) and *Olea*. Both species, together with others that are less abundant in the palaeobotanical record, like *Arbutus unedo*, *Laurus nobilis* and *Rhamnus*, could have formed part of cantabric holm oak-type communities (Table 3). Dispersed formations are still found on the northern Spanish coast, surviving in areas with poor dry soils, normally limestone areas where little water is retained. It is a relict community that continues to survive owing to the mild climate on the Atlantic coast (Bueno and Fernández, 1991; Meaza, 1991), and which has been documented in the palaeobotanical record since at least the Late Glacial period (Ramil et al., 1998). It is evident between ca. 8300 and 5300 cal BC at sites like Urdaibai (Iriarte et al., 2006), Playaundi (Sanchez-Goñi, 1996), Kobeaga II (Iriarte, 2000) and Pareko Landa (Iriarte et al., 2007/2008). At sites in inland valleys in the Atlantic zone the vegetation dynamic is similar, with the difference that further from the coast, the marine influence is reduced and thermophilic-sclerophilic elements like holm oak are scarcer. The predominance of deciduous formations is shown above all by the spread of hazel copses in the pollen record, together with other deciduous species like alder, birch, deciduous *Quercus*, linden

and walnut, as shown by the studies carried out at sites like Arenaza (Isturiz and Sanchez-Goni, 1990), Ekain (Dupré, 1984), Urtiaga (Sanchez-Goni, 1993) and Sustriği (Iriarte et al., 2007/2008), between ca. 10,600 and 7500 cal BC. The same trend is recorded in some nearby marine sequences, where deciduous woodlands were dominant in the Early Holocene (Naughton et al., 2007; da Silva, 2012).

Table 3

Summary of the main vegetation features of the Holocene in the Western Pyrenean Region.

Chronocultural phases	Main vegetation features
Epipaleolithic-Mesolithic (ca. 9500–5500 cal BC)	
Pyrenees	- In the high altitudes preponderance of conifer formations (<i>Pinus sylvestris</i>) - In lower altitudes dominance of deciduous woodlands (<i>Corylus</i>, <i>Betula</i>, <i>Alnus</i>, deciduous <i>Quercus</i>, <i>Fraxinus</i>, <i>Salix</i>, <i>Populus</i>).
Athlantic area	- Spread of deciduous forests (deciduous <i>Quercus</i>, <i>Corylus</i>, <i>Ulmus</i>, <i>Alnus</i>, <i>Betula</i>, <i>Tilia</i>). - Lou presence of <i>Pinus sylvestris</i>. - In the coastal area presence of thermophilous copses, cantabric holm oak-type communities with <i>Quercus ilex</i>, <i>Olea</i>, <i>Arbutus unedo</i>, <i>Laurus nobilis</i> and <i>Rhamnus</i>.
Transition area	- Dominance of deciduous woodlands, likely shaping mixed forest (<i>Corylus</i>, <i>Betula</i>, <i>Alnus</i>, deciduous <i>Quercus</i>, <i>Fraxinus</i>, <i>Salix</i>, <i>Populus</i>). - The low representation of Pine (<i>Pinus sylvestris</i> <10%) indicates its presence at a regional scale.
Early Neolithic (ca. 5500–4500 cal BC)	
Pyrenees	- Predominance of deciduous woodlands (<i>Corylus</i>, <i>Betula</i>, deciduous <i>Quercus</i>, <i>Tilia</i>, <i>Alnus</i>, <i>Ulmus</i>) and pines (<i>Pinus</i> sp.). - First emergence of anthropic palynomorphs (<i>Plantago</i> sp.). - Increasing presence of shrubs.
Athlantic area	- Predominance of deciduous woodlands (deciduous <i>Quercus</i>, <i>Corylus</i>, <i>Ulmus</i>, <i>Alnus</i>, <i>Betula</i> and <i>Tilia</i>). - First emergence of anthropic palynomorphs (<i>Plantago</i> sp.). - First evidences of cultivated plants (cereal pollen and cultivated seeds).
Transition area	- Predominance of deciduous woodlands (<i>Corylus</i>, deciduous <i>Quercus</i>, <i>Ulmus</i>, <i>Alnus</i>). - Increasing open areas. - Emergence of pastures of anthropogenic origin (<i>Chenopodiaceae</i>, <i>Urtica dioica</i> type, <i>Plantago lanceolata</i>, <i>Sporormiella</i> sp. and <i>Sordaria</i> sp.).
Middle-Late Neolithic and Copper Age (ca. 4500–2200 cal BC)	
Pyrenees	- Predominance of mesophilous forests with <i>Corylus</i>, deciduous <i>Quercus</i>, <i>Betula</i>, <i>Ulmus</i>, <i>Alnus</i>, <i>Tilia</i> and <i>Fraxinus</i>). - Regression in values of arboreal pollen. - Emergence of crop fields (cereal pollen grains).
Athlantic area	- Prevalence of deciduous arboreal vegetation (<i>Corylus</i>, deciduous <i>Quercus</i>, <i>Ulmus</i>, <i>Alnus</i>, <i>Fraxinus</i>). - Development of open areas. - Emergence of crop fields (cereal pollen and cultivated seeds).
Transition area	- Dominance of deciduous forests (<i>Corylus</i>, deciduous <i>Quercus</i>, <i>Ulmus</i>, <i>Alnus</i>, and <i>Fraxinus</i>). - From the Chalcolithic increase of evergreen forests in the southern sector (<i>Quercus ilex</i>, <i>Buxus</i>). - Regression in values of arboreal pollen. - Cereal pollen >3% and cultivated seeds.
Bronze Age and Iron Age (ca. 2200–200 cal BC)	
Pyrenees	- Presence of deciduous forests (deciduous <i>Quercus</i>, <i>Corylus</i>, <i>Alnus</i>, <i>Betula</i>) and conifers (<i>Pinus sylvestris</i>). - Spread of <i>Fagus</i>. - Agro-pastoral activities.
Athlantic area	- Dominance of deciduous forests (deciduous <i>Quercus</i>, <i>Corylus</i>, <i>Tilia</i>, <i>Fraxinus</i>, <i>Ulmus</i>). - Spread of <i>Fagus</i>. - Crop fields and anthropogenic pastures.
Transition area	- Presence of deciduous forests and development of evergreen forest in the southernmost area (evergreen <i>Quercus</i>, <i>Pinus sylvestris</i>, <i>Buxus</i>). - Spread of <i>Fagus</i>. - Crop fields and anthropogenic pastures

In the Pyrenees, palynological studies of both archaeological deposits and peat bogs display the characteristic arboreal colonisation in the early Holocene. However, in this area, pine forests predominate, together with smaller deciduous stands formed by hazel, birch, alder and deciduous *Quercus*. This is seen at the archaeological sites of Abauntz, Zatoya and Berroberría (López-García, 1982; Boyer-Klein, 1984) between 10,000 and 6300 cal BC. The montane environment clearly explains the preponderance of conifer formations (pine forests) over deciduous woodlands, which are of secondary importance at the start of the Holocene. In other deposits located at lower altitudes, such as the sites of Aizpea and Padre Areso, once more deciduous woodland is the main vegetation type. In this case, hazel is the main element in the Mesolithic forest during the whole seventh millennium cal BC (Iriarte, 2001; Iriarte et al., 2007/2008). It is accompanied by other typical trees in

mixed deciduous woodland (birch, alder, deciduous *Quercus*, ash, willow, poplar), as well as some pines and, in Padre Areso, holm oaks (*Q. ilex* type), owing to its more southern location (Table 3).

In a more inland area, in the climate Transition zone, several archaeological sites are known in a relatively small radius (Sub-Mediterranean valleys), all of the same chronology. As may be expected, they all reflect a very similar environmental dynamic in

which the area was fully occupied by deciduous woodland, a mixed forest where hazel predominated over deciduous oak, birch, lime, yew and willows. Pine is present in the record but with low percentages (<10%), possibly indicating its presence at a regional scale. This is seen in the palynological studies at Mendandia (Iriarte, 2006), Kanpanoste (Sánchez-Goni, 2004) and Kanpanoste Goikoa (Iriarte, 1998) from ca. 6800 to 6000 cal BC.

4.2.2. Early Neolithic (ca. 5500–4500 cal BC). First evidence of farming

In the sixth millennium cal BC, a new archaeological reality can be recognised in the Iberian Peninsula: the Neolithic technocomplex. It appears in a cultural process, within which a series of transformations resulted in the Mesolithic way of life (hunter-gatherers) changing to another situation in which food was

produced by agriculture and animal husbandry. The new economic practices also led to a series of social transformations which are equally reflected in the vegetation, mainly because of the need for open spaces for the new productive activities.

The Western Pyrenees has traditionally been considered a marginal area as regards the adoption of Neolithic ways of life, supposedly with a delay in the arrival of the first Neolithic evidence (Zapata et al., 2004; Zapata and Peña-Chocarro, 2012). However, several research projects in the last decade have questioned this view, as they have documented certain novelties associated with the Neolithic technocomplex dated in the sixth millennium cal BC at such sites as the rock-shelter of Peña Larga (5700–5500 cal BC) (Fernández-Eraso, 2007/2008), Cueva de Marizulo (5500–5200 cal BC) (Alday and Mujika, 1999), Atxoste (5300–5000 cal BC) (Alday et al., 2012), Cueva de Arenaza (5200–4800 cal BC) (Arias and Altuna, 1999) and the open-air settlement of Herriko Barra (5200–4700 cal BC) (Iriarte et al., 2004). One particularity of the Neolithic in this area is the clear existence of two types of archaeological sites in terms of the inhabitants' forms of subsistence. At one type, intense use of wild resources has been documented (Abauntz, Aizpea, Herriko Barra, Kanpanoste Goikoa, Pico Ramos, Zatoya), whereas the other type exhibits marked evidence of pastoral or agricultural activity (Arenaza, Kobaederra, Peña Larga). This reveals a wide range of situations, possibly derived from the origin and evolution of the early Neolithic in the region.

In the Atlantic zone, several palaeoenvironmental studies have noted the first evidence of anthropisation of the landscape in the late sixth millennium and during the fifth millennium cal BC, such as the deposits of Pico Ramos ca. 4800–4500 cal BC (Zapata, 2002) and Pareko Landa (Iriarte et al., 2007/2008), where deforestation was incipient. However, in these deposits no evidence of farming practices has been documented, unlike at other nearby sites, such as Herriko Barra ca. 5200–4700 cal BC (Iriarte et al., 2004), where cereal pollen was identified. This is one of the earliest evidence of the development of agriculture on the northern coast of Iberia, at a site where neither pottery nor domestic animals were recorded and

which, therefore, appears to have specialised in the use of wild resources in a diet complemented by cereals. Other palaeobotanical remains of cereals, in this case grains, were identified at some neighbouring sites, like Kobaederra ca. 4700–4300 cal BC (Zapata, 2002). Unlike at the coastal sites of Pico Ramos and Herriko Barra, the identification of domestic fauna at Kobaederra seems to indicate an economic focus on domestic resources, especially livestock but also agriculture.

In the Pyrenees, studies at such sites as Abauntz (López-García, 1982), Zatoya (Boyer-Klein, 1989) and Aizpea (Iriarte, 2001) suggest the predominance of deciduous woodland during the early Neolithic (hazel, birch, deciduous *Quercus*, linden, alder). The same trend is documented in some peat bogs (Table 3). In the case of the new palynological study of Atxuri presented here, the diagram shows at the PAZ ATX-1 (ca. 5568–3434 cal BC) a predominance of hazel, reaching 50% of terrestrial pollen, and other deciduous woods, with *Quercus*, birch and elm as main components, together with pines (Fig. 2). Although discreet, some indicators of incipient anthropisation have been detected in this area, such as preponderance of hazel, and a residual presence of anthropic palynomorphs and heaths (López-Sáez et al., 2007/2008). However, no farming practices have been detected in any of these records, which can undoubtedly be related to the continuation of Mesolithic ways of life in these mountain environments, where hunting and gathering took priority. Some evidence of anthropisation is also recognisable in peat bogs. At Belate, in the first half of the fifth millennium cal BC (ca. 5000–4500 cal BC), heliophilous and anthropic taxa and shrubs proliferated, attesting some disturbance to the natural woodland vegetation (Peñalba, 1989).

In the climate Transition zone, the oldest references correspond to the site of Mendandia. Palynological data (Iriarte, 2006) document a steep decline in arboreal pollen in the central part of the sixth millennium cal BC (ca. 5700–5300 cal BC) in comparison with the Mesolithic, as it only reaches a percentage of about 9%. Hazel is still the most common tree (Table 3), together with other deciduous species, like oaks, lime and others found in very small proportions (alder, walnut and elm). This significant decline in the

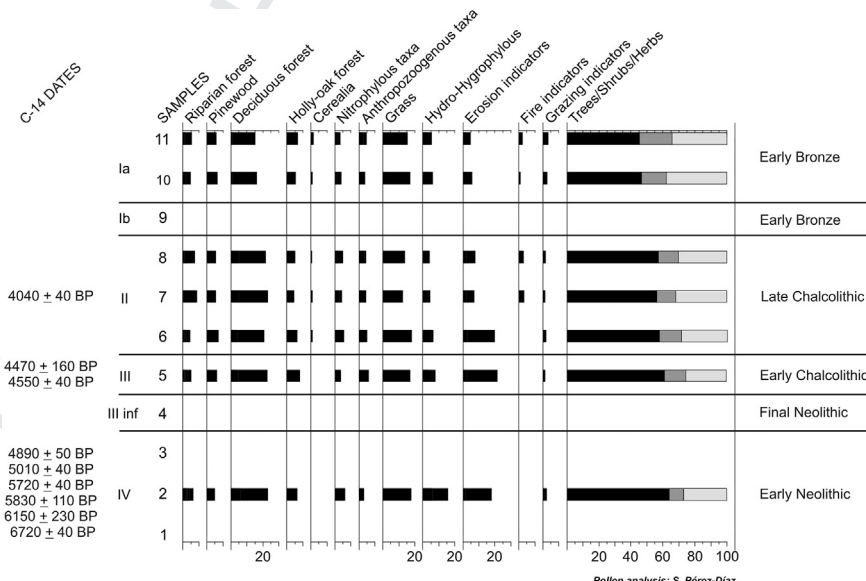


Fig. 3. Palynological synthetic histogram of the archaeological site of Peña Larga (Pérez-Díaz, 2012). Riparian forest (*Alnus*, *Fraxinus*, *Salix*, *Ulmus*), Pinewood (*Pinus sylvestris* type, *Pinus pinaster*), Deciduous forest (*Corylus*, *Betula*, deciduous *Quercus*, *Tilia*), Holly-oak forest (evergreen *Quercus*), Cerealia (*Cerealia* type), Nitrophylous taxa (*Centaurea nigra* type, *Dipsacus fullonum* type, *Aster* type, *Cardueae*, *Cichorioideae*, *Convolvulus arvensis* type), Anthropozoogenous taxa (*Chenopodiaceae*, *Plantago* sp.), Grass (*Poaceae*), Hydro-hygrophylous (*Cyperaceae*, *Filicales* Monolete, *Polypodium vulgare* type), Erosion indicators (*Glomus* cf. *fasciculatum*), Fire indicators (*Chaetomium* sp.), grazing indicators (*Sordaria* sp., *Sporormiella* sp.).

representation of forest masses is undoubtedly connected with human activity, as attested by a large increase in synanthropic pollen taxa, above all Compositae liguliflorae, and to a lesser extent Dipsacaceae and Compositae tubuliflorae, which often accompany human presence.

In the southern mountains, in Sierra de Cantabria, the early Neolithic in the deposit of Peña Larga, ca. 5700–4500 cal BC (Fernández-Eraso, 2007/2008) is characterised by the existence of deciduous woodland in the surroundings of the rock-shelter, in which hazel was the most common species, together with deciduous *Quercus*, birch, linden, etc. (Fig. 3). Economic practices appear to have been focused on pastoralism (Pérez-Díaz, 2012), according to the palynological evidence, with the characteristic vegetation of nitrified soils (anthropozoogenous pastures with *Chenopodiaceae*, *Urtica dioica* type, *Plantago lanceolata*), as well as coprophilous fungi (*Sporormiella* sp. and *Sordaria* sp.) (van Geel, 1976, 1978, 2006; van Geel et al., 1983, 2003; López-Sáez et al., 2000). These conclusions are also supported by the archaeozoological study which identified remains of domestic cattle, ovicaprines and pigs, as well as some wild animals (Castaños, 1997). The first results of DNA analysis of osseous remains point in the same direction (Fernández-Eraso, 2010). In the immediate vicinity, the rock-shelter of Los Husos I has yielded cereal pollen in ca. 5200–4800 cal BC (Fernández-Eraso, 2007/2008).

4.2.3. Middle-Late Neolithic and Copper age (ca. 4500–2200 cal BC). Consolidation of the producing economy

The appearance of a broad series of novelties associated with the Neolithic technocomplex, from the sixth millennium cal BC onwards, marked the start of change in the ways of life of the communities that had lived in the region until then. These changes affected vital aspects of human society, such as the structuring of space, settlement patterns, social relationships, funerary behaviour, and so on. They become particularly clear in the Middle-Late Neolithic (ca. 4500–3200 cal BC) and the Chalcolithic (ca. 3200–2200 cal BC). Additionally, they left an unquestionable impact on the vegetation.

In the Atlantic zone, palynological studies show the prevalence of deciduous arboreal vegetation, typical of temperate and humid conditions (Table 3). Mixed forests predominated, above all with hazel and oak, together with Rosaceae, elm, ash, alder, etc. However, this woodland displayed clear evidence of anthropisation, mainly in the form of human deforestation, with the presence of nitrophilous herbaceous taxa between ca. 4200 and 2500 cal BC, in Katillotxu (Iriarte, 2007/2009) and Pico Ramos (Iriarte, 1994b). The latter site also yielded data about the adoption of agriculture in this part of the coastal strip, as the carpological study cites remains of *Hordeum vulgare* (together with wild plants) in the late fifth and early fourth millennia cal BC (4229–3798 cal BC) (Zapata, 2002). Deforestation is generally recorded in most of the archaeological deposits, but is less apparent in natural deposits, like in Saldropo (Peñalba, 1989). In these cases, the scant alteration of the vegetation due to anthropic action may be explained by the absence of archaeological sites in their vicinity, which meant they were relatively unaffected by prehistoric activity.

In the Pyrenean zone, several deposits are known. The same general pattern is seen in them all: the dominance of mesophylous woods (hazel, oak). In the case of Atxuri (last part of PAZ ATX-1, PAZ ATX-2 and beginning of PAZ ATX-3), other deciduous taxa appear at this moment, such as alder and ash, beginning their continuous presence in the sequence (Fig. 2). Diagrams also show the interaction of human communities with all Pyrenean ecosystems, on a much larger scale than in the Early Neolithic. Some very clear signs of anthropisation have been noted in all the pollen sequences. Accordingly, the earliest evidence of agriculture in the Navarran

Pyrenees dates back to the fourth millennium cal BC, as cereal pollen has been identified ca. 4000–3500 cal BC in Quinto Real (Galop et al., 2001), ca. 3250 cal BC in Artxilondo (Galop, 2005), and ca. 2300 cal BC in Atxuri.

In the Transition zone, deposits with palynological studies reflect the vegetation consisting of deciduous woodland, dominated by hazel and deciduous *Quercus*, together a significant percentage of riparian vegetation (elm, alder, ash), which indicates a high level of humidity. However, this vegetation was noticeably affected by anthropic deforestation already begun in the Early Neolithic (Table 3). The driving force behind this process of deforestation was farming activity. In this period, most deposits have yielded cereal pollen in percentages of over 3%, which is the value usually taken as a sign of crops being grown in the immediate surroundings of the deposit (Diot, 1992; López-Sáez and López-Merino, 2005). In the same way, the diagrams reflect several indicators of livestock in the area around the site. As well as nitrophilous herbaceous communities favoured by the presence of animals, a non-pollen micro-fossil of coprophile ecology (*Sordaria* sp.) has been recorded, a good indicator of livestock *in situ* (van Geel et al., 1983, 2003). In Sierra de Cantabria, a large number of archaeological sites (Peña Larga, San Cristóbal, Los Husos I and Los Husos II) all show that in the Late Neolithic, the predominant vegetation was still temperate and humid woodland, with deciduous trees like hazel, birch, oak (possibly Portuguese oak) and lime above all, forming a mixed deciduous forest. However, from the climate viewpoint, certain change to more arid conditions is shown in the Chalcolithic by plant species better adapted to warm dry climates, such as box, and a certain preponderance of holm oak formations (Fig. 3). This intense vegetation change was forced by the farming economy. During the Chalcolithic, agriculture became widespread across the region, after the first introductory phase. In several deposits with earlier levels, this seems to be the moment when agriculture developed. Another economic activity visible in the palynological record is animal husbandry. The identification of anthropozoogenic flora, forming part of pastures grazed by livestock, the presence of coprophilous fungi, micromorphological evidence of accumulations of dung, and archaeozoological data draw a picture in which pastoralism was one of the main activities of Middle-Late Neolithic and Chalcolithic communities (Pérez-Díaz, 2012).

4.2.4. Bronze Age and Iron Age (ca. 2200–200 cal BC). The rise of complex societies

Our understanding of the last prehistoric societies, defined in Europe as Bronze Age and Iron Age communities, has advanced a great deal in recent decades, with interesting changes from social, cultural and economic perspectives identified (Fernández-Eraso et al., 2009; Sesma et al., 2009). These changes indicate growing socioeconomic complexity, and are reflected clearly in the vegetation (Carozza et al., 2005; Pérez-Díaz and López-Sáez, 2012).

In the Atlantic zone, numerous deposits of this time are known, of various kinds including caves, rock-shelters, open-air sites and megalithic monuments, as well as natural deposits. They all exhibit the predominance of woodland consisting mainly of deciduous *Quercus*, hazel, linden, ash and elm, as well as the start of the beech expansion (Table 3). These forests display a high level of anthropic alteration. This is most evident in studies carried out at archaeological sites and less so in natural deposits. The palynological record from Saldropo peat bog, in a transitional biogeographic environment between the Atlantic and Mediterranean sides of the watershed, ca. 2200–1700, reflects characteristics noticeably different from those documented at archaeological sites (Peñalba, 1989). Whereas in the latter the main trait is intense anthropisation and the preponderance of vegetation with an anthropic origin, at

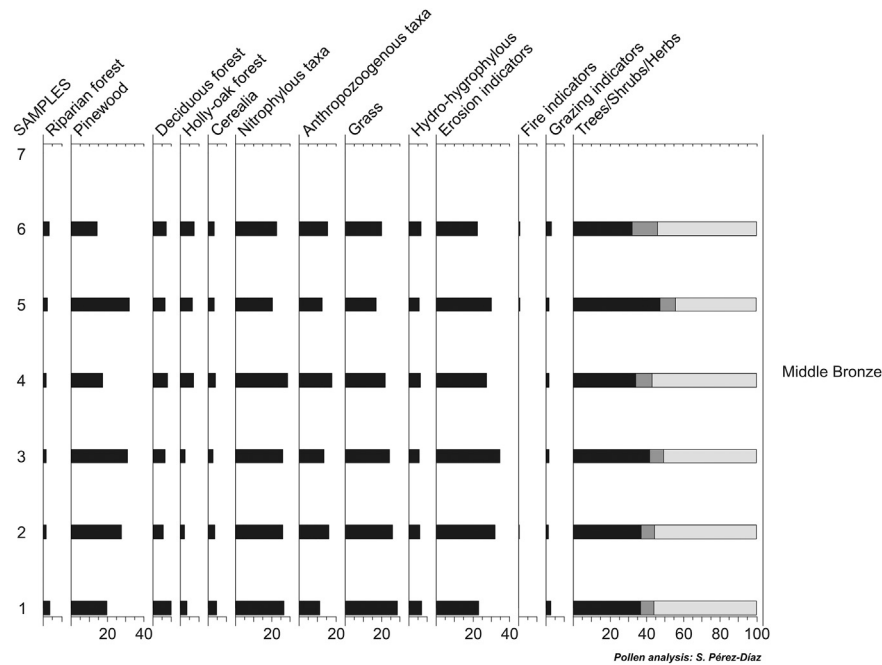


Fig. 4. Palynological synthetic histogram of the archaeological site of Masparra (Pérez-Díaz and López-Sáez, 2013). Riparian forest (*Alnus*, *Salix*), Pinewood (*Pinus sylvestris* type, *Pinus pinaster*), Deciduous forest (*Corylus*, *Betula*, deciduous *Quercus*, *Tilia*), Holly-oak forest (evergreen *Quercus*), Cerealia (Cerealia type), Nitrophyllous taxa (*Centaurea nigra* type, *Dipsacus fullonum* type, *Aster* type, *Cardueae*, *Cichorioideae*), Anthropozoogenous taxa (*Chenopodiaceae*, *Plantago* sp., *Polygonum aviculare* type, *Urtica dioica* type), Grass (*Poaceae*), Hydro-hygrophyllous (*Filicales Monoete*, *Filicales Trilete*, *Polypodium vulgare* type), Erosion indicators (*Glomus* cf. *fasciculatum*, *Pseudoschizaea circula*), Fire indicators (*Chaetomium* sp.), Grazing indicators (*Sordaria* sp., *Sporormiella* sp.).

Saldropo hardly any signs of alteration to the vegetation can be detected. Arboreal cover is important, mainly of deciduous taxa like *Quercus* and *Corylus*. The start of the beech expansion is dated at the same time. Similarly, the scarcity of synanthropic elements supports the hypothesis of little human activity in that area.

In the Pyrenees, some deposits are known that cover the last two millennia cal BC. In the cave of Zatoya, the palynological study (Boyer-Klein, 1989) reflected an increase in the representation of pine forests. The vegetation inferred from palynological studies of the peat bogs at Belate (Peñalba, 1989), Atxuri (Fig. 2) and Quinto Real (Galop et al., 2001, 2004) show a similar general dynamic. In the case of Atxuri, presented herein, the diagram shows the development of a mesophyllous forest, as previously, dominated by deciduous *Quercus* and hazel. At the same time, the expansion of beech woods is identified; from ca. 2500 cal BC it detects its continuous presence in the environment and from ca. 1700 cal BC a drastic increases. From this moment beech woods becomes the dominant forest. Certain anthropic presence is detected by the presence of cultivated plants and nitrophyllous communities (Fig. 2).

In the climate Transition zone, tree cover is generally reduced in the surroundings of human settlements and consists mainly of deciduous elements, although in the southernmost areas the advance of herbaceous formations better adapted to dry environments can be detected, in addition to the predominance of pine and holm oak among the trees and box among the bushes (Table 3). Farming activity seems to be the driving force behind the intensification of anthropic influences, which is seen in pollen diagrams in the existence of areas with no arboreal vegetation. Practically all the settlements with palynological studies display high levels of cereal pollen, which indicates the relative proximity of agricultural land. In the same way, the identification of herbaceous communities typical of pastures, together with the presence of ascospores of coprophilous fungi, attest the extent of livestock farming and environmental pressure (Fig. 4). This is very clear both in valley

sites like Ameztutxo, Masparra, Mendizabala and Los Castros de Lastra, and in deposits in mountain areas, like those at over 900 m altitude in Sierra de Cantabria (Peña Larga, Peña Parda, San Cristóbal) (Fig. 3).

5. Conclusions

The evolution of the landscape during the Holocene in the Western Pyrenees is the result of the interaction of many factors, giving as final result the current landscape configuration. Among these factors we must mention the increasing human activity, especially evident from the first moments of the Neolithic. Moreover, the climatic evolution in this environment has contributed in the evolution of plant communities.

The combination of archaeobotanical studies, specially palynological, in both archaeological sites and natural deposits is a very interesting tool for the reconstruction of vegetation history and the dynamics of anthropisation both spatially and temporally, particularly the development of farming activity (agriculture and pastoralism) and their impact on the environment (deforestation, erosion, fires, etc.).

After the Late Glacial period, characterised by cold and dry conditions, a more benign climate developed in the early Holocene, with increasing humidity and higher temperatures. These new conditions favoured arboreal expansion, and consequently a reduction in the area occupied by the steppe herbaceous communities. The predominant vegetation type in the early Holocene was mixed woodland with deciduous species like hazel, oak, birch, alder, lime, etc., better adapted to the new environmental conditions. Pine forests have also been recorded, especially in high mountain areas of the Pyrenees and more continental conditions in the south of the region.

The first changes associated with the new cultural technocomplex of the Neolithic are seen in the Western Pyrenees at an

early date, in the late sixth millennium cal BC. At that time, palynological studies have documented animal domestication (presence of ruderal and nitrophilous vegetation and coprophilous fungi) and agriculture (cereals) in several deposits. The new farming way of life in the Neolithic is the reason for the first evidence of anthropisation of the landscape, such as deforestation and plant species with an anthropic origin. Vegetation in the Early Neolithic (ca. 5500–4500 cal BC) is characterised by an advance in herbaceous formations with an anthropic origin and a still significant presence of mesophyllous forest with hazel, oak, birch, lime, elm, walnut, etc. The economic practices of the Early Neolithic (agriculture and livestock farming) are not documented together in the pollen record at all sites, as in some places the use of wild resources co-exists with pastoralism or agriculture, attesting a wide range of situations possibly deriving from the origins and evolution of the Early Neolithic in the region.

During the Late Neolithic (ca. 4500–3200 cal BC) and the Chalcolithic (ca. 3200–2200 cal BC) the farming economy is consolidated, as shown by the progressive intensification of anthropisation within the spread of cereal-growing and animal husbandry. The forests, that had previously dominated the vegetation were profoundly affected by the phenomenon of sedentarisation and the increasing importance of Neolithic and Chalcolithic farming activity, with the result that areas colonised by herbaceous communities with an anthropic origin became more common. The dominant vegetation type is characteristic of temperate humid conditions, with deciduous forests of hazel, oak, birch, alder, etc. However, in the southern part of the region (to the south of Sierra de Cantabria), vegetation adapted to a drier and warmer climate is detected, with a predominance of holm oak and the increasing importance of box.

Prehistoric communities in the Bronze Age (ca. 2200–900 cal BC) and the Iron Age (ca. 900–200 cal BC) lived in a mostly deforested environment due to farming activity. Both cereal-growing and livestock farming are widely documented in all the deposits and formed the economic bases of these groups. This landscape exhibited deciduous woodland greatly reduced in size, as a consequence of continuing and increasing anthropisation, with some pines in mountain areas and holm oak on poorer soils and warmer areas. However, the landscape was dominated by pastures with an anthropic origin and others with an anthropozoogenic vocation.

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