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PII: S1040-6182(20)30149-X

DOI: https://doi.org/10.1016/j.quaint.2020.03.042

Reference: JQI 8211

To appear in: Quaternary International

Received Date: 31 January 2020

Revised Date: 21 March 2020

Accepted Date: 22 March 2020

Please cite this article as: Irabien, Marí.Jesú., Cearreta, A., Gómez-Arozamena, José., Gardoki, J., Martín-Consuegra, Aitor.Ferná., Recent coastal anthropogenic impact recorded in the Basque mud patch (southern Bay of Biscay shelf), *Quaternary International* (2020), doi: https://doi.org/10.1016/j.quaint.2020.03.042.

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	Journal 116-proof
1	RECENT COASTAL ANTHROPOGENIC IMPACT RECORDED IN THE BASQUE
2	MUD PATCH (SOUTHERN BAY OF BISCAY SHELF)
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14	Abstract
15	The historical anthropogenic impact on sediments from the Basque Mud Patch
16	(southern Bay of Biscay) is explored using a multidisciplinary approach including the
17	analysis of natural (²¹⁰ Pb) and artificial (¹³⁷ Cs, ^{239/240} Pu) radiotracers, major elements
18	(Al, Mn), metals (Zn, Pb, Cu, Cr), Pb isotopic ratios, and foraminiferal and pollen
19	contents. The study of three short cores (19-46 cm), despite being hindered by the
20	effects of biomixing, allow the calculation of a sedimentation rate of 1 ± 0.1 mm yr ⁻¹ .
21	Distribution with depth of Al-normalised concentrations of metals reflects an increasing
22	trend since 1880 CE, related to the industrialization of the Basque coastal area.
23	According to the Sediment Quality Guidelines applied, contents of Zn and Pb appear
24	as a potential cause of concern, given that they exceed the values from which adverse
25	biological effects can be occasionally expected. However, foraminiferal assemblages
26	do not show recognizable changes along the cores following increasing trace metal
27	concentrations. Finally, pollen results reveal an increasing trend of coniferous taxa and

28 a parallel reduction of authochthonous Deciduous Quercus, probably as a

consequence of reforestation works. Data obtained confirm that effects of coastal
anthropogenic activities extend to the adjacent shelf, where muddy deposits are likely
to act as a trap for contaminants.

32

33 Keywords: Metals; Foraminifera; Radiotracers; Pollen; Mud depocenter; Bay of Biscay

34

35 1. Introduction

36 The world's continental shelves, which underlie for little more than 8% of the global 37 marine surface, are key interfaces between terrestrial sediment source areas and 38 deep-sea depositional systems (Covault and Fildani, 2014). They are complex 39 environments characterised by the interplay among many factors, such as tectonic 40 setting, material inputs from landmasses, climate, hydrodynamic processes (tidal flows, 41 waves, currents) and human activities (Nittrouer et al., 2007; Chiocci and Chivas, 42 2014). These marine areas can act as significant sinks not only for fluvial and estuarine 43 materials, but also for anthropogenic contaminants and carbon compounds (Mahiques 44 et al., 2016; Qiao et al., 2017), which are frequently associated with fine-grained 45 particles (Förstner, 2002). Many clastic shelves host mud deposits (the so-called 46 depocenters), showing different three-dimensional geometries and characteristics, 47 where land-supplied fine-grained materials are retained (Hanebuth at al., 2015). In the 48 Bay of Biscay (Fig. 1), mud deposits appear as isolated patches (Hanebuth at al., 49 2015), conditioned by the favourable morphological context, short-term events of high 50 fluvial sediment discharge and oceanic current forcing factors (Lesueur et al., 2002; 51 Jouanneau et al., 2008). Mud patches located on the French shelf have been largely 52 investigated in previous works (Castaing and Allen, 1981; Jouanneau et al., 1989; 53 Lesueur et al., 2001, 2002; Schmidt et al., 2005; Dubrulle et al., 2007; Mengual et al., 54 2019; Mojtahid et al., 2019), whereas that lying along the northern Iberian margin 55 adjacent to the Basque coast (Fig. 1) has received much less attention (see references

56 below). For this reason, the study of these fine-grained deposits represents the main57 goal of this work.

58 The Basque region is a relatively small area (7,234 km²) located in the innermost 59 part of the Bay of Biscay (Fig. 1). Its recent history has been decisively marked by the 60 exploitation and processing of local iron ores by steel industry, upon which a sizeable 61 chunk of the economic growth of the city of Bilbao and the whole region was based for 62 over a century. Unfortunately, these activities were carried out within a historical 63 framework of unsustainable development, and almost all river catchments, estuaries 64 and coastal water bodies have been affected, to a different extent, by a wide array of 65 anthropogenic pressures, including urban, mining and industrial discharges, 66 hydromorphological modifications, dredgings, and disposal of solid wastes (Tueros et al., 2009; Borja et al., 2016). Consequences have been especially dramatic in coastal 67 68 areas where, in the current scenario of climate change, changes due to sea-level rise 69 are likely to be overwhelmed by human impacts (Chust et al., 2009).

70 In the Basque Country, people living in coastal areas (around 1,500,000 inhabitants) 71 accounts for 70% of the total population (IHOBE, 2017). About 55% settles in two 72 metropolitan zones around the cities of Bilbao and San Sebastian-Donostia (Fig. 1) 73 that make up less than 10% of the whole Basque territory. Most original transitional 74 domains have been lost as a result of land reclamation and all estuaries show, to a 75 greater or lesser extent, evidences of sediment contamination (Cearreta et al., 2004; 76 Legorburu et al., 2013b; Irabien et al., 2015). Increased concentrations of trace metals 77 and PAHs appear also in nearshore samples (Belzunce et al., 2004; Borja et al., 2008; 78 Tueros et al., 2009; Viñas et al., 2010; Legorburu et al., 2013a), but enrichment levels 79 are in general distinctly lower than those determined within the estuaries. Evidences of 80 human influence are likely to extend to deeper areas, given the high contents of priority 81 and emerging micropollutants determined in the Capbreton canyon, on the French part 82 of the Bay of Biscay (Boutier et al., 2011; Azaroff et al., 2020).

This is the first work to assess the potential impact left by human activities in the muddy deposits of the Basque continental shelf. For this purpose, three sediment cores were studied using a multiproxy approach that includes the analysis of radionuclides (²¹⁰Pb, ¹³⁷Cs and ²³⁹⁺²⁴⁰Pu), trace metals, isotopic signatures of Pb, foraminiferal assemblages and pollen contents.

88

89 2. Regional setting and background

90 The continental shelf adjacent to the Basque Country (Fig. 1) is a narrow area (7-20 91 km) about 150 km long which receives the discharge of suspended material (around 1.57x10⁶ t yr⁻¹) of twelve different river systems (Pascual et al., 2004; Uriarte et al., 92 93 2004). As early observed by De Buen (1933), coarse to medium-sized sandy materials 94 occur in the western part, while in the eastern section the mid- and outer shelves (50-95 <150 m depth) are covered by an extended unit of fine-grained sediments (Rey and 96 Mendialdea, 1989; Jouanneau et al., 2008; Galparsoro et al., 2010), henceforth 97 expressed as the Basque mud patch (BMP) (Fig. 1). It is bounded by rocky outcrops in 98 a shoreward direction and extends to the shelf-break and the edges of the Capbreton Canyon. According to Jouanneau et al. (2008), this deposit is about 680 km² in size 99 100 and has a maximum thickness of 7 m.

101 The study of the benthic microfauna in surface sediments from the BMP developed 102 by Pascual et al. (2008) evidenced that distribution of foraminifera and ostracod 103 assemblages were mainly controlled by environmental variables such as calcium 104 carbonate, particulate organic carbon and silt-clay proportion. These authors found 105 specimens affected by the oil-spill derived from the sinking of the tanker MV Prestige, 106 which introduced about 63,000 metric tonnes of heavy fuel in the Bay of Biscay 107 between November 2002 and February 2003 (Laffon et al., 2006). In addition, data 108 provided by Fernández Martín-Consuegra (2018) from the analysis of both surface and 109 cored sediment samples pointed out the role of cool-waters upwelling through the San 110 Sebastian Canyon to explain foraminiferal distribution during the last 4000 years.

111

112 3. Materials and methods

113 3.1. Sampling

Sediment cores were obtained during the EUSKASED (Euska2) scientific cruise of the R/V *Côtes de la Manche* in July 2004 from the central area of the BMP, in front of the city of San Sebastian-Donostia (Fig. 1). Cores KS-04 (43°25.45N, 2°0898W, 44 cm long) and KI-03 (43°25.69N, 2°09.18W, 19 cm long) were retrieved at 135 m depth using a Kullenbeg corer, whereas core KI-06 (43°25.424N, 1°57.747W, 46 cm long) was obtained at 119 m depth using an interface corer (Champilou, 2005). Sediment cores were cut in 1 cm-thick intervals at the University of Bordeaux (France).

121 3.2. Radiometric analysis

Specific activities of both natural (²¹⁰Pb y ²²⁶Ra) and artificial (¹³⁷Cs, ^{239/240}Pu and 122 123 ²³⁸Pu) radiotracers were measured in samples from core KS-04. In addition, activity concentrations of other natural radionuclides such as ²³²Th, ²¹²Pb and ²²⁸Ac, which 124 125 occur in the disintegration chain of ²³⁵U, and ⁴⁰K were also determined. Except ^{239/240}Pu and ²³⁸Pu, they were analysed in the University of Cantabria (Spain) by gamma 126 127 spectrometry, using a low-background high purity HPGe detector (see procedural 128 details in Alvarez-Iglesias et al., 2007 and Cearreta et al., 2013). Sediment samples 129 were homogenised, hermetically sealed and stored for at least 25 days to ensure secular equilibrium between ²²⁶Ra, ²²²Rn and the short-lived daughter nuclides of the 130 latter. Excess ²¹⁰Pb activities (²¹⁰Pb_{xs}) were determined by subtracting the ²²⁶Ra activity 131 (assumed to equal the supported ²¹⁰Pb activity) from the total ²¹⁰Pb activity. Measures 132 of ^{239/240}Pu and ²³⁸Pu were carried out in the Laboratory of Environmental Radioactivity 133 134 of the University of Valencia (Spain) using alpha spectrometry, after a sequential 135 chromatographic separation process (analytical details are described in Sáez-Muñoz et 136 al., 2020).

Dating methods using ²¹⁰Pb_{xs} are based on the assumption that the tracer cycle is a 137 closed system, so that the flux of ²¹⁰Pb_{xs} to the sediments is constant and there is no 138 139 post-depositional remobilization (Goldberg, 1963; Krishnaswami et al., 1971). In 140 coastal and marine environments, which are considered as open systems (Pfitzner et 141 al., 2004), flux of ²¹⁰Pb to the sediments is mainly due to direct atmospheric deposition, 142 but contributions can also be produced by advection, the flow of suspended matter from river discharges and the disintegration of ²²⁶Ra dissolved in the water column, 143 144 although this last process is likely to be significant only at great depths (Cochran et al., 145 1990). These factors may introduce certain limitations in the application of the method 146 as well as in the different models for dating. However, and despite these uncertainties, 147 ²¹⁰Pb activities along the studied core showed a clearly exponential decrease with 148 depth, reflecting that accumulation rates were constant. Therefore, both sediment 149 accumulation rates (SAR) and mass accumulation rates (MAR) were estimated through 150 the Constant Flux (CF:CS) model (Koide et al., 1973; Robbins, 1978), which assumes constant ²¹⁰Pb flux and sedimentation rates. 151

152 3.3. Geochemical analysis

153 Sediment samples from all cores were mechanically homogenised using an agate 154 mortar and pestle. Elemental concentrations were analysed in Activation Laboratories 155 Ltd. (Actlabs, Ontario, Canada). An amount of 0.5 gr of sample was digested with agua 156 regia (a combination of concentrated hydrochloric and nitric acids) for 2 hours at 95°C. 157 Samples were then cooled, diluted with deionized water and analysed using a Varian ICP. Detection limits were 0.01% for AI, 5 mg kg⁻¹ for Mn, 2 mg kg⁻¹ for Zn and Pb, and 158 159 1 mg kg⁻¹ for Cu and Cr. Quality Control for the digestion was 15% for each batch, 2 160 method reagent blanks, 6 in-house controls, 8 sample duplicates and 5 certified 161 reference materials. An additional 20% CQ was performed as part of the instrumental 162 analysis to ensure quality in the areas of the instrumental drift.

When dealing with sediments, it is an extended practice to normalize trace metal concentrations to the contents of a conservative element such as aluminium to compensate for the potential effects of granulometric and mineralogical differences (e.g. Mil-Homens et al., 2006; Ho et al., 2012). Sediments along each core are fairly homogeneous, and vertical profiles displayed by measured concentrations and Alnormalised values for metals are remarkably similar. However, only Al-normalised concentrations are represented to facilitate inter-core comparisons.

Lead isotope analyses were carried out at the Geochronology and Isotope Geochemistry-SGIker facility of the University of the Basque Country UPV/EHU (Spain) using a Thermo Fisher Scientific MC-ICP-MS Neptune, after digestion of the samples with Aqua Regia. NBS981 standard was used for calibration and accuracy check measurements.

175 3.4. Microfaunal study

176 Benthic foraminiferal analysis was performed for a 1-cm sample every two, from 0-1 177 cm to 16-17 cm depth in core KI-03 and from 0-1 cm to 18-19 cm depth in core KI-06. 178 Samples were sieved through a 63 µm mesh and washed with tap water to remove 179 clay- and silt-size (mud) fractions. They were dried at 40 °C and preliminarily observed 180 under a stereoscopic binocular microscope using reflected light. All contained a great 181 abundance of well-preserved foraminifera and tests were picked until a representative 182 amount of >300 individuals was obtained for each sample. In total, 19 samples and 183 >6,500 foraminiferal tests were studied and taxonomically classified following Loeblich 184 and Tappan (1988) updated in the World Register of Marine Species (WoRMS).

185 3.5. Pollen analysis

Samples from core KI-03 were analysed at each 1-cm layer every two from 1-2 cm to 17-18 cm depth (alternating with those of foraminifera) and in core KI-06 pollen samples coincided with those of foraminifera. They were chemically (HCI, KOH, HF) and densimetrically (Thoulet solution) treated for the extraction of pollen and non-

pollen palynomorphs (NPPs) at the Institute of History-CSIC in Madrid. In addition, each sample was mounted in glycerine for the correct observation and identification of palynomorphs under an optical microscope at 400x and 600x magnification. Taxonomic classification was performed following Goeury and de Beaulieu (1979), Moore et al. (1991), Reille (1992) and Cugny et al. (2010). A minimum of 500 pollen grains were counted per sample while aquatic taxa, fungal spores and other NPPs were excluded from the terrestrial pollen sum (Robles-López et al., 2018).

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198 4. Results and discussion

Ranges of Mn and trace metal concentrations in the three studied cores andregional background values proposed by different authors are included in Table 1.

201 4.1. Core KS-04

202 Sediments from this core are fine-grained (mean granulometry ranges from 14.8 to 203 26.6 μ m), exhibit rather constant contents of CaCO₃ (15.8-23.6%), and X-ray 204 radiograph reveals undisturbed horizontal lamination, pointing out low bioturbation intensity (Champilou, 2005). Moreover, concentrations of ²³²Th, ²¹²Pb, ²²⁸Ac and ⁴⁰K 205 206 display almost vertical profiles, pointing out the absence of significant compositional 207 changes along the core. Unfortunately, samples from the topmost 3 cm and those 208 corresponding to 12-13 cm and 16-18 cm depth were not available for analysis. 209 Therefore, the lack of the uppermost sediments prevents processes of near surface 210 biomixing to be discarded. In any case, the exponential decrease with depth observed for the excess activity of ²¹⁰Pb (Fig. 2) allows, using the slope of the line derived from 211 the linear regression between In²¹⁰Pb_{xs} and depth (Appleby and Oldfield, 1992) a mean 212 213 sedimentation rate of 1 ± 0.1 mm yr⁻¹ to be calculated. If some mixing were occurring 214 throughout the core, the estimated sedimentation rate would be biased high (Silverberg 215 et al., 1986). However, the obtained value is similar to the lowest rates determined by 216 Jouanneau et al. (2008) in the sedimentary deposits of the BMP (between 1.3 and 5.0

mm yr⁻¹) and by Schmidt et al. (2005) in the adjacent mud depocenters of the French
shelf (from nearly 1 mm yr⁻¹ up to almost 3 mm yr⁻¹).

Transient tracers, such as ¹³⁷Cs and to a much lesser extent ^{239/240}Pu, have been 219 220 widely applied in dating sediments, usually to corroborate sedimentation rates 221 determined using ²¹⁰Pb. In undisturbed cores, the first occurrence of both radioisotopes 222 should correspond to the beginning of atmospheric nuclear weapons testing in the 223 early 1950s, whereas highest values are usually attributed to the maximum global fallout in 1962-64 CE. In this core, ¹³⁷Cs and significant amounts of ^{239/240}Pu appear for 224 225 the first time at 6 cm depth (Fig. 2), slightly deeper (1 cm) than expected according to 226 the ²¹⁰Pb-derived ages and yielding a comparable but somewhat higher sedimentation rate (~1.3 mm yr⁻¹). However, ²¹⁰Pb-based chronological model is likely to provide a 227 228 better resolution for the maximum concentrations of both radioisotopes found at around 229 4 cm depth (mid-1960s) than that derived from the first appearance of ¹³⁷Cs (which 230 should indicate that they were deposited later in time, in the 1970s), given that they 231 probably reflect the peak flux preceding the ban of atmospheric testing of nuclear 232 weapons (year 1963). The anomalous penetration of radiotracers to greater depths 233 than expected, previously described in several works (Livingston and Bowen, 1979; 234 Jaakkola et al., 1983; Cochran, 1985; Crusius et al., 2004) confirms that sediment 235 accumulation may not be the only factor controlling their distribution in sediments 236 (Moon, 2003). Biomixing is probably the most powerful mechanism for displacement, being ¹³⁷Cs (and virtually every other contaminant of interest) a lot more sensitive to 237 mixing than ²¹⁰Pb (Johannessen and Macdonald, 2012). Moreover, vertical profiles of 238 ¹³⁷Cs can be also distorted by other phenomena such as desorption from sediment 239 240 grains and diffusive transport (Hancock et al., 2011) or vertical displacement of pore-241 waters caused by changes in density of the overlying column (Holby and Evans, 1996). 242 Vertical profiles displayed by Al-normalised concentrations of Zn, Pb and Cu are 243 very similar in shape, showing low and almost constant values below 13 cm depth and 244 a progressive increase towards more recent sediments (Fig. 3). The chronological

245 framework built on the basis of the ²¹⁰Pb-derived sedimentation rate indicates that 246 these sediments were deposited before 1880 CE. Although the Basques' knowledge 247 on iron mining and forging probably antedated Roman times (Douglass and Bilbao, 248 2005), the change to the modern steel industry began in 1848 with the launch of the 249 first blast furnace near the city of Bilbao. However, the main expansion took place in 250 the 1880s owing to the implementation of new technologies that enabled the 251 conversion of local low-phosphorous iron ores to high-quality steel (Bilbao, 1983). 252 Therefore, concentrations of Zn, Pb and Cu in these bottom sediments seem to 253 represent well pre-industrial levels, whereas those deposited above reflect the 254 increasing impact of recent anthropogenic activities. This is supported by the vertical 255 evolution of Pb isotopic ratios, which exhibit a significant shift towards depleted values 256 as concentrations of this metal increase (Fig. 4a). While samples deposited before 257 1880 CE appear together within the field associated with pre-industrial estuarine 258 sediments, thereafter there is an increasing input of sources less radiogenic than the 259 background component (Fig. 4b). For comparison, potential mixing end-members have 260 been plotted (Fig. 4b), and they include local galenas from the Arditurri district (see Fig. 261 1 for location), coal, metallurgical emissions, and leaded gasolines. Obtained data tend 262 to fall along a linear array with tighter adherence for the greater ratios, representative 263 of background values, and a greater spread at the smaller ratios, indicative of the 264 increasing contribution of anthropogenic sources (most probably metallurgical works) 265 since 1880 CE.

With regard to Cr, when compared to Zn, Pb and Cu, it can be observed that the onset of clearly increased contents is located upwards in the sedimentary column (Fig. 3). This is consistent with the results obtained in sediment cores collected from the Bilbao estuary, main industrialised area of the Basque Country, which indicated that Cr and Ni entered into the estuarine domains later in time (Cearreta et al., 2002). Finally, it is noteworthy that values of Mn/Al with depth mimic the behaviour observed for the other trace metals, showing a gradual increase above 10 cm depth (Fig. 3). Mn is a

273 redox-sensitive element that has been extensively used to test potential postdepositional remobilization in sediment cores, given that it can be solubilized in the 274 275 anoxic environment and redeposited in the upper oxic layers (Valette-Silver, 1993; 276 McKay et al., 2007; Kuzyk et al., 2017). In fact, previous studies developed on 277 sediments from the French sector of the Bay of Biscay have identified marked peaks of 278 this element in the upper section of the cores (at depths between 0 and 5 cm) related 279 to early diagenetic processes, with maximum contents located at the oxygen 280 penetration depth (Hyacinthe et al., 2001; Mouret et al., 2009; Boutier et al., 2011). 281 Notwithstanding this, in this core potential diagenetic effects are likely to be masked by 282 anthropogenic inputs of Mn (frequently used in steelmaking), as enhanced 283 concentrations of this element (accompanied by high contents of other pollutants such 284 as Zn, Pb or Cu) have been detected in coastal and estuarine sediments from the 285 Basque Country (Ramos et al., 1990; Tueros et al., 2009; Larreta et al., 2012; 286 Legorburu et al., 2013a; Rodríguez-Iruretagoiena et al., 2016; Garmendia et al., 2019), 287 particularly in areas of slag disposal adjacent to the Bilbao estuary (Borja et al., 2008).

Assuming the constant flux of ²¹⁰Pb_{xs} and the CF:CS model, a constant mass 288 accumulation rate of 97 \pm 6 mg cm⁻² yr⁻¹ was estimated. Calculation of the metal fluxes 289 290 for sediments deposited before 1880 CE, which are supposed to represent pre-291 industrial contributions, vielded the following values (in mg m⁻² yr⁻¹): 407 for Zn, 116 for 292 Pb, 68 for Cu, 136 for Cr and 1,510 for Mn. These values are within the ranges 293 proposed by Álvarez-Vázquez et al. (2017) as natural fluxes for estuarine sediments of 294 Galician Rias (northwestern Spain). However, the intense industrialization of the 295 Basque coastal area is likely to have resulted in an increase of the metal fluxes to the 296 BMP, since they are currently around 6 times higher for Pb, 4 for Zn, 3.5 for Cu, 2 for 297 Mn and 1.5 for Cr.

4.2. Cores KI-03 and KI-06

Sediments from cores KI-03 and KI-06 are composed by fine-grained materials
 (13.7-19.2 μm and 14.1-16.6 μm respectively) and exhibit almost constant contents of

301 CaCO₃ (19.6-23.0% and 17.2-19.7% respectively) and evidences of bioturbation 302 (deformed horizontal lamination) (Champilou, 2005).

303 In both cores distribution patterns of Al-normalised concentrations of Mn and trace 304 metals with depth bear striking similarities (Fig. 5). They show low values in bottom 305 sediments, which increase irregularly upwards until reaching constant levels in the 306 upper layer (about 10 cm thick). These flattened uppercore profiles are likely to reflect 307 not only the anthropogenic inputs of metals during the last decades, but also the 308 homogenization effect exerted by biomixing. The only exception to this trend is Mn, 309 which exhibits enhanced contents in surface samples. This enrichment could be 310 related to the diagenetic remobilization of this element from reducing subsurface 311 sediment layers and the subsequent precipitation upwards under oxic conditions. 312 Moreover, in both cores there is a marked horizon at 16 cm depth with elevated Al-313 normalised concentrations of Mn, Zn, Cu and Pb, which can reflect both an 314 anthropogenic origin (an episode of contamination) or a post-depositional effect 315 (migration of the redox front and scavenging of metals by Fe-Mn oxyhydroxides). 316 Therefore, further studies involving pore waters and metal speciation should be made 317 in this area to gain insight into the Mn redox cycling, as it can lead to remobilization of 318 elements within the sedimentary column and trigger/accelerate the transfer of trace 319 elements from sea water (Tribovillard et al., 2006).

320 A total number of 81 different foraminiferal species were found in both cores. The 321 relative abundance of porcellaneous (average 3.4%, range 0.6-6.9%) and agglutinated 322 tests (1.6%, 0.3-4.6%) was minor compared to dominant hyaline forms (95.0%, 90.9-323 99.1%). The average number of species per sample was 37 (range 22-54) and those 324 taxa present more abundantly in all samples were Bolivina difformis (11.2%, 4.5-325 19.5%), Bolivina spathulata (10.6%, 2.3-19.9%), Globocassidulina subglobosa (9.1%, 326 2.8-20.4%), Bolivina pseudoplicata (7.6%, 3.8-10.5%), Rosalina irregularis (7.1%, 3.2-327 16.1%), Gavelinopsis praegeri (7.1%, 2.7-10.9%), Cibicidoides lobatulus (6.9%, 3.6-328 11.8%), Cassidulina laevigata (5.3%, 1.9-10.3%), Bulimina marginata (4.4%, 3.2-5.7%)

and *Hyalinea balthica* (2.3%, 0.3-5.4%) (Fig. 6). This assemblage is made mainly of
species tolerant to low oxygen conditions and does not exhibit a significant variation
through time along the cores. Pascual et al. (2008) and Martínez García (2012) found
these same taxa as dominant in the surface samples of this muddy area.

333 Forty different pollen taxa were observed in both cores. Tree pollen shows the 334 greatest dominance (83.5%, range 63.9-96.6%) whereas herbaceous (10.7%, 2.8-335 23.2%) and shrub (5.8%, 0.6-14.9%) pollen are secondary (Fig. 7). Hydro-hygrophytes 336 (27.4%, 9.4-47.1%) and NPPs (1.8%, 0.3-4%) were also present in the samples but not 337 included into the terrestrial pollen sum. The most significant taxa along both cores were 338 Pinus pinaster (63.3%, 34.7-84.8%), Cyperaceae (6.3%, 1-14.5%), Deciduous Quercus 339 (6.3%, 1.9-14.1%), Pinus sylvestris type (5.5%, 1.5-13.2%), Apiaceae (4.5%, 0.4-340 13.9%), Erica type (3.0%, 0.2-8.5%), Betula (2.6%, 0.8-6.5%), Alnus (2.2%, 0.6-5.9%), 341 Poaceae (1.8%, 0.4-5.7%), Arbutus type (1.6%, 0.2-4.6%) and Cichorioideae (1.5%, 342 0.2-3.6%). There is a progressive increase in coniferous species (P. pinaster and P. 343 sylvestris type) with time and a parallel reduction of Deciduous Quercus, herbaceous 344 (Apiaceae and Poaceae) and shrub (Arbutus type and Erica type) taxa. It should be 345 noted the testimonial appearance of Eucalyptus in some of the most superficial 346 samples, a clearly exotic taxon introduced since mid-20th century for economic 347 reasons.

348 4.3. Environmental assessment

349 According to Ridgway and Shimmield (2002), for most temperate zones the 350 evidence suggests that the impacts of contaminated sediments on shelf seas tend to 351 be small and largely confined to the coastal area adjacent to the estuaries, given that 352 contaminants present in the suspended sediment load are rapidly spread and diluted 353 by mixing with less contaminated marine sediment. However, results obtained in the 354 studied cores reflect an increasing trend in metal concentrations. Table 1 provides an 355 overview of the different background values proposed for coastal sediments of the 356 Basque Country (Legorburu et al., 1989; Cearreta et al., 2002; Rodríguez et al., 2006)

357 and the North-East Atlantic (OSPAR, 2006). It is well known that factors such as 358 particle size, mineral composition and organic carbon content of the samples, as well 359 as the analytical method used, can condition to a great extent the estimation of pre-360 industrial references and probably explain the observed differences. Therefore, in this 361 work the consistently low concentrations found in downcore samples have been 362 considered to represent pre-industrial conditions and, in order to provide a means to 363 assess the relative role of anthropogenic impact on the BMP, enrichment factors have 364 been calculated as follows (Herut and Sandler, 2006):

365 Enrichment factor (EF): [X(s)/Al(s)] / [X(b)/Al(b)]

366 where X is metal concentration, (s) is sample, and (b) is background value.

367 Ranges of metal concentrations in the three cores are very similar (Table 1), with 368 the highest EFs oscillating between 1.3 and 1.4 for Cr, 1.9-2.2 for Mn, 2.5-3.2 for Cu, 369 2.9-3.8 for Zn and 4.7-5.8 for Pb. However, it is noticeable that, contrary to that 370 observed in the historically polluted sediments from some Basque estuaries (Ramos et 371 al, 1990; Cearreta et al., 2018; Irabien et al., 2018; 2019), metal contents in samples 372 from the BMP do not exceed in any case the effects range-median (ERM) values 373 proposed by Long et al. (1995) as sediment quality guidelines (see Table 1), which 374 represent conditions where adverse biological effects occur frequently. In fact, levels of 375 Cu and Cr in all samples remain even below the effect range-low (ERL) values, which 376 correspond to concentrations below which adverse biological effects are rarely 377 observed. Based upon this evaluation, Pb and Zn appear as main contaminants of 378 environmental concern in this area, given that they exhibit concentrations above the 379 ERLs and below the ERMs until 6 cm depth in core KS-04 and until 10 cm depth in 380 cores KI-03 and KI-06. Therefore, they enter the "possible-effects" range, within which 381 adverse biological effects would occasionally occur. In fact, foraminiferal assemblages 382 do not show any significant variation through time in the cores despite trace metal 383 concentrations increase towards more recent times. This situation was previously 384 observed in the recent sedimentary record of the Bilbao estuary, where Cearreta et al.

(2002) concluded that occurrence of foraminifera in the recent industrial zones and
 along the modern estuary was not related to defined levels of trace metals.

387 As expected, concentrations of metals in the BMP are distinctly lower than those 388 determined in some Basque estuaries (Belzunce et al., 2004; Montero et al., 2013), 389 where high amounts of contaminated sediments remain accumulated (Cearreta et al., 390 2002; 2018). Nevertheless, an overall improvement in estuarine sediment quality has 391 been observed over the last years as a consequence of the decrease of contaminant 392 disposal activities and the implementation of efficient systems for wastewater treatment 393 (Leorri et al., 2008; Tueros et al., 2009; Legorburu et al., 2013b). In historically polluted 394 areas such as the Bilbao estuary, the study of sediment cores indicates that the 395 gradual deposition of "cleaner" sediments is leading to a marked decline in trace metal 396 concentrations towards recent samples (Irabien et al., 2018; 2019). Unfortunately, no 397 similar trends can be recognised in cores from the BMP (Figs. 3 and 5).

Regarding the dynamics of the vegetation, the progressive increase of coniferous taxa is indicative of their extensive plantations carried out on the Cantabrian coastal area since the end of the 19th century resulting in the continuous decrease of authorhonous Deciduous *Quercus*. This was exacerbated from the 1940s, when the "National Reforestation Plan" was launched (Michel Rodríguez, 2006).

403

404 5. Conclusions

405 This work confirms that the influence of the anthropogenic activities developed in 406 the Basque coast since late 19th century extends to the marine environment, where 407 muddy deposits (BMP) act as a trap for the chemicals released. The similarity between 408 the results obtained in the three studied cores seems to reflect that Mn and trace 409 metals are widely distributed over the BMP. As expected, due to the mixing with 410 uncontaminated marine sediments, concentrations are significantly lower than those 411 determined historically in the nearby estuaries, which are the main source of sediments 412 to the shelf. However, and contrary to that determined in these estuarine areas, no

413 significant improvement is observed in recent times, probably as a consequence of the 414 low sedimentation rates and the disturbing effects of biomixing processes. Therefore, 415 confirming the extent of contamination within marine muddy deposits and monitoring its 416 evolution in the future appear as essential tasks for the environmental assessment and 417 responsible management of continental shelves adjacent to industrialised areas.

418

419 Acknowledgements: All samples were kindly provided by Dr Ana Pascual (UPV/EHU). 420 Aintzane Goffard (UPV/EHU) prepared samples and produced foraminiferal results 421 from core KI-06. This research was funded by Spanish MINECO (CGL2013-41083-P 422 and RTI2018-095678-B-C21, MCIU/AEI/FEDER, UE), UPV/EHU (UFI11/09) and 423 EJ/GV (IT976-16) projects. Aitor Fernández Martín-Consuegra was supported by a 424 predoctoral grant from the Basque Government (PRE_2017_1_0173). The authors 425 thank technical and human support provided by SGIker (UPV/EHU/ERDT, EU). Two 426 anonymous reviewers improved the original manuscript with their comments and 427 constructive suggestions. This is contribution 53 of the Geo-Q Zentroa Research Unit 428 (Joaquín Gómez de Llarena Laboratory).

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774	FIGURE CAPTIONS
775	
776	Fig. 1. a) Simplified map of the Bay of Biscay with the main continental margins
777	(modified from Martínez-García, 2012). b) Sedimentary map of the Basque shelf
778	(modified from Jouanneau et al., 2008) and location of sampling sites (white dots) in
779	the Basque Mud Patch.
780	
781	Fig. 2. Distribution of ²¹⁰ Pb _{xs} , ¹³⁷ Cs and ^{239/240} Pu (Bq kg ⁻¹) with depth in core KS-04.
782	
783	Fig. 3. Distribution of Al-normalised concentrations of metals with depth in core KS-04.
784	
785	Fig. 4. a) Dispersion plot of concentrations of Pb versus ²⁰⁸ Pb/ ²⁰⁷ Pb ratios in samples
786	from core KS-04. b) Dispersion plot for ²⁰⁶ Pb/ ²⁰⁷ Pb versus ²⁰⁸ Pb/ ²⁰⁷ Pb ratios. Full circles
787	represent sediments from the BMP (this work): grey circles indicate samples deposited
788	before 1880 CE and black ones those deposited later. Full triangles and empty circles
789	are used for pre-industrial samples from Basque (Leorri et al., 2014) and French
790	estuaries (Elbaz-Poulichet et al., 1984; Monna et al., 1997) respectively. Squares
791	represent galenas from the Arditurri district (Velasco et al., 1996). Empty inverted
792	triangles indicate Spanish coals (Alvarez-Iglesias et al., 2007; Díaz-Somoano et al.,
793	2007), whereas diamonds were used for metallurgical emissions from Northwestern
794	France (Véron et al., 1999). Full inverted triangles represent leaded gasolines used in
795	France and the United Kingdom (Monna et al., 1997) and in Spain (Alvarez-Iglesias et
796	al., 2007).
797	
798	Fig. 5. Distribution of Al-normalised concentrations of metals with depth in cores KI-03
799	and KI-06.
800	
801	Fig. 6. Relative abundance (%) of dominant foraminiferal species with depth in cores
802	KI-03 (above) and KI-06 (below).
803	
804	Fig. 7. Palynological diagram of cores KI-03 (above) and KI-06 (below) including the
805	relative abundance of the most representative pollen taxa and NPPs. Black dots
806	indicate the presence of <i>Eucaliptus</i> in the upper samples.
807	

Table 1. Ranges of metal concentrations in the analysed cores, regional pre-industrial values for coastal areas proposed by different authors, mean contents determined in downcore samples (n=12) and Sediment Quality Guidelines used in this work (all in mg kg⁻¹).

	Mn	Pb	Zn	Cu	Cr
KS-04	145-323	9-74	30-168	7-26	13-20
KI-03	162-303	15-95	49-163	10-24	16-22
KI-06	167-349	10-89	51-175	10-27	17-27
Background values					
Cearreta et al. (2002)	400	23	60	18	76
Legorburu et al. (1989)	289	15.8	56.1	20.9	25.6
Rodríguez et al. (2006)	240	31	174	33	26
OSPAR (2006)	-	25	90	20	60
Downcore samples (this work)	154	14	52	9	16
Sediment Quality Guidelines					
ERL (Long et al., 1995)		46.7	150	34	81
ERM (Long et al., 1995)		218	410	270	370















Declaration of interests

 \boxtimes The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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